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HARVARD COLLEGE OBSERVATORY

AND THE

NUMERICAL ANALYSIS LABORATORY

OF THE MASSACHUSETTS INSTITUTE OF TECHNOLOGY

TECHNICAL REPORT NUMBER TEN

A COMPARATIVE ANALYSIS OF ATMOSPHERIC
DENSITIES FROM METEOR DECELERATIONS
OBSERVED IN MASSACHUSETTS AND NEW MEXICO

BT

LUIGI G. JACCHIA

CAMBRIDGE, 1952

HARVARD COLLEGE OBSERVATORY AND THE NUMERICAL ANALYSIS LABORATORY OF THE MASSACHUSETTS INSTITUTE OF TECHNOLOGY

Technical Report Number Ten

Work performed under Contract DA-19-020-ORD-1093 for the Office of Ordnance Research, U.S. Army

A COMPARATIVE ANALYSIS OF ATMOSPHERIC DENSITIES FROM METEOR DECELERATIONS OBSERVED IN MASSACHUSETTS AND NEW MEXICO

by

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Cambridge, 1952

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FOREWORD

The meteor photographs which form the basic material of this report were taken at the Meteor Stations of the Harvard College Observatory under Task D of NOrd 8555 and Task I of NOrd 10449 for the U.S. Naval Bureau of Ordnance, and now under NSori-U7647 for the Office of Naval Research. The measurement and reduction of the plates and the analysis of the resulting data was done at the Numerical Analysis Laboratory (formerly Center of Analysia) of the Massachusetts Institute of Technology under Tasks D and I of NOrd 8555 and of NOrd 10455 and Task A of NOrd 10455 for the U.S. Naval Bureau of Ordnance, and now under DA-19-020-ORD 1093 of the Office of Ordnance Research, U.S. Army.

The persons involved in the project were:

a) At the Harvard Stations:

Harlow Shapley, Director of the Harvard College Observatory Fred L. Whipple, Project Director

Frances W. Wright, in charge of plate inspection and records

Richard E. McCrosky, observer-in-charge at the New Mexico Stations.

Following is a list of the observers at the New Mexico Stationa; the number in parenthesia indicates the number of months served prior to January 1951 (date of the last plate included in this report):

Philip S. Carrol! (29), Peter O. Cioffi (12), Keith Quard (2),
Ivvin Levitan (4), Gunther Schwartz (6), Harlan J. Smith (4),
Henry J. Smith (2).
Plate inspectors in New Mexico: Catherine Carroll (29), Juanita Engle (7).

b) At the M. I.T. Numerical Analysis Laboratory:

Zdenék Kopal, Director of the M.I.T. Numerical Analysia Laboratory and Project Director until June 1951.

Luigi G. Jacchia, Project Supervisor until June 1951; Director of the M.I.T. Numerical Analysis Laboratory and Project Director after July 1951.

Computera: Virginis K. Brenton, Robert E. Briggs, Jeannie R. B. Carmichael, Francia G. Davoren, Diana H. Mason, Laurelle B. Parrotta, Dorothy T. Pemberton, Louise Richardson, Martha B. Shapley.

Occasional Computers: Mary H. Baker, David D. Brown, Carolyn S. Littlejohn, Helen M. Pillana, Anita Porell, Sidney Shapiro, Charlotte G. Treuenfela.

Assistant Computers: Heater L. Bassett, Helen M. Carr, Belle J. Helpern.

Most of the computations involved in the present report were performed by Mr. Briggs, Miss Carmichael and Miss Pemberton.

Much of the information concerning the first fifteen Massachusetts moteora (Table II) was supplied by Dr. F. L. Whipple, who had reduced them before the inception of the present project.

ABSTRACT

Velocities, accelerations and integrated luminosities were determined on photographic plates for 46 Massachusetts meteors and 73 New Mexico meteors. Atmospheric densities, computed from the New Mexico meteors in the same general fashion as in Technical Report No. 2, are in good agreement with the slope of the density profile derived from V-2 rockets; however, in analogy with the previous results for Massachusetts meteors, the residuals from the rocket curve are strongly dependent on velocity. A change of one full power in v in the fundamental equation is needed to make the residuals independent of velocity, but then the slope of the density curve diverges considerably from the rocket profile. The introduction of an empirical scale factor R reduces the velocity-free densities back again to the rocket profile. These final densities show only a slight seasonal effect, if any.

Massachusetts observations were reduced using the same constants as for New Mexico. The final densities are in good agreement with the New Mexico densities up to 75 km, but above that height the Massachusetts densities are systematically higher. The seasonal effect is more than twice as large for the Massachusetts observations as for New Mexico. The bad distribution of the Massachusetts meteors with regard to seasons, velocities and heights should, however, suggest some caution in interpreting this last result.

1. The Observational Material:

Only double-station meteors are included in the present report. All photographs through August 11, 1948, were taken in Massachusetts, from the Cambridge and Oak Ridge (now Agassiz) stations, for which co-ordinates are given below:

Station	Oak Ridge (Agassiz)	Cambridge
λ (Greenw.)	71°33′29″.82	71° 7'45".45
ϕ	+42°30′20″.72	+42°22′53″.70
h	190.2 m.	18.3 m.

All photographs from August 14, 1949, inclusive, were taken in New Mexico, from the twin stations near Las Cruces, whose co-ordinates are:

Station	Doña Ana	Soledad Canyon
λ (Greenw.)	106°47′58″.50	106°36′42″.32
φ	+ 32°30′21″.94	+ 32°18′13″.61
h	1412.3 m.	1567.4 m.

The baseline between the Massachusetts stations is 37.896 km long; the baseline in New Mexico measures 28.567 km.

The cameras with which the meteors were photographed are listed in Table I.

TABLE I

Meteor Cameras

a) Massachusetta

	Oak Rid	ge (Age	ussiz)			C	-b rid	ge	
Camera	Ap.	f	r.p.m.	n	Comera	Ap.	f	r.p.m.	n
AI	1.5	6	600	2	FA	1.5	6	600	2
KB	3.0	7	1200	2	KA	3.0	7	1800	2
• AC	1.5	13	0	0	KL	1.5	6	630.5	4
• MC	16	83	0	0					

b) New Mexico

	Doi	na Ana				Soled	lad Car	iyon	
Camera	Ap.	f	r.p.m.	n	Camera	Ap.	f	r.p.m.	n
ΑI	1.5	6	740	4	FA	1.5	6	630	4
KB	3.0	7	1800	2	KA	3.0	7	1200	2
KF	3.0	7	1800	2	KE	3.0	7	1200	2
KH	3.0	7	0	0	KG	3.0	7	0	0

Ap. and f are the aperture and the focal length, in inches; r.p.m. is the number of revolutions per minute of the rotating shutter associated with the camera, and n is the number of measurable features (breaks and/or dots) per shutter revolution. The cameras marked with an asterisk are not regular meteor cameras, but other astronomical instruments on which the meteor was photographed by accident.

Basic data for all meteors are given in Tables II and III. The Massachusetts material includes 63 meteors. Of these, 46 could be used to determine velocities and deceleration and yielded 86 atmospheric densities. The New Mexico material comprised 83 meteors, 73 of which yielded 168 density values.

The following is an explanation of those columns in Tables II and III which are not self-explanatory. Q is the angle of intersection on the celestial aphere of the two meteor trails, $Z_{\rm R}$ is the zenith distance of the spparent radiant, $v_{\rm co}$ and $u_{\rm co}$ are the extrapolated velocity (in km/sec) and mass (in grams) of the meteor before it entered the earth's atmosphere; t is the duration of the meteor as photographed from station A (Oak Ridge in Massachusetta, Dona Ana in New Mexico), $H_{\rm vin}$ is the magnitude, reduced to the visual scale, which the meteor stained at maximum light.

In the "shower" column all those meteors which do not belong to well-recognized showers were listed as sporadic. There are, however, two new showers in the list:

- 1) The Virginids, based on three meteors in 1939, 1942 and 1950, all between March 18 and March 21, with geocentric velocities of about 31 km/sec. Approximate radiant: $\alpha = 185^{\circ}$, $\delta = +55^{\circ}(1950.0)$.
- 2) The x Cygnids, based on four 1950 meteors between August 19 and August 22; geocentric velocity 27 km/sec, radiant: α = 292°, δ = +55°(1950.0).

The fundamental data used in the computation of atmospheric densities are given in Tables IVa and IVb for Massachusetts and Va and Vb for New Mexico meteors. In these tables Fr indicates the fraction of the tropic year, ϕ the parameter $360^{\circ} \times Fr$, ν the velocity in km/sec, $d\nu/dt$ the acceleration in km/sec², p.s. the probable error of the acceleration; the three last quantities were taken at the center of s least-squares solution involving n shutter breaks (form of the equation of condition: $D = a + bt + ce^{kt}$; D = distance of break in space from a

fixed point, t = time, k = a pre-determined constant). H is the height above sea level, in km, and m the mass of the meteor in grams at the time for which ν and $d\nu/dt$ were taken. The meaning of ρ_1 , ρ_2 , ρ_3 and ρ_4 is explained in section 2; p and w are the weights computed according to equations (3) and (4), respectively. Δ_1 , Δ_2 , Δ_3 , Δ_4 are the residuels of $\log \rho_2$, $\log \rho_3$ and $\log \rho_4$ from profile C (Table VI); Δ_3' and Δ_4' (for Massachusetts) are residuals of $\log \rho_3$ and $\log \rho_4$ from profile M.

Data pertaining to heights up to 90 km are to be found in Tables IVa and Va; Tables Va and Vb give the equivalent information for heights greater than 85 km. Data for heights between 85 and 90 km can thus be found in both seta of tables, liated in different fashion. Although for the computation of the mean densities the data for H>85 km were handled in groups rather than individually, some of the analyses required individual atmospheric densities up to 90 km; this explains the overlap between the tables.

Meteor No. 2278 (New Mexico, 1950 Dec. 11.2) deserves a special mention. This meteor akismed the upper atmosphere with an angle of incidence of only 12° and a velocity of 59 km/sec. Its photographic length is 53°; its first half was caught by the AI and FA cameras, while the accord half was recorded by the KE camera. On both the AI and the FA plates the trail is extremely long and extends from a point not far from the center clear to the edge of the field. The acceleration computed from both plates is small, but positive, and in our opinion this is due to the progressive change in the images of comparison stars and breaks from the center to the edge. The large number of breaks on these exceptionally long trails makes these spurious accelerations such larger than their probable errors and this, in turn, would lead to unduly high weights for the resulting (negative) atmospheric densities, with the danger of vitiating the analysis. In view of this situation, it was deemed necessary to eliminate the AI and FA densities for this meteor from the present investigation.

2. Atmospheric Densities from Meteor Data; Weights and Weighted Means:

In a first approximation the atmospheric densities ρ were computed in the same fashion as in Technical Report No. $2^{(1)}$, using the formula

$$\rho = -K \frac{1}{a^3} v^{-2} \frac{dv}{dt} . \tag{1}$$

The mass m was computed by equation (9) of Technical Report No. 2, which shall be repeated here for easy reference:

$$= \frac{2}{\tau_0} \int_{t}^{t} \frac{I}{t^{3}} dt . \tag{2}$$

 τ_0 is the luminous-efficiency coefficient and I the visual intensity of the acteor (4), derived from the photographic magnitude with the correction +1.8.° For a detailed description of the photometric methods followed, see Technical Report No. 3⁽²⁾, pages 5-11. The densities thus computed, using K=1, are designated as ρ_1 .

Each individually determined acceleration was used to obtain a value of ρ_1 . Accelerations determined from different plates for the same meteor were treated independently of each other. If one photographic trail yielded two or more decelerations, these were also handled separately.

The basic quantities which enter into the computation of ρ_1 are affected by observational errors in varying degrees. While the velocity v is always known to an accuracy of 1% or better, the acceleration dv/dt may have probable errors ranging from 0.01 to 10 times it a actual value, and the reliability of the probable error itself will vary according to the number of ahutter breaks used to compute the acceleration. The mass π , which is a function of the integrated light intensity, has an element of uncartainty in the extrapolation to zero of the intensity curve at the very end of the trajectory. This uncertainty is reflected in a

^{*}Twenty-five New Nexico meteors for which both visual and photographic magnitudes were evallable confirm this correction. Their mean is ± 1.89 .

ABLE II

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8	71.	8	8.	ઙ	18	8	\$	8.	 E.	•.15	3	10.	6	•.15	ð.	-18	.31	3	ä	.13	. 25	=-	2	•.16	8	33	ä	8.	×	9:	8	• 16	ķ	.30
108 03	5.71	.s.	S.2	8.8	6.46	6.28	6.40	7.9	6.7	6.22	6.9	6.51	6.61	6.6	90.7	4.5	8.7	7.16	71.17	6.83	8.7	6.88	2.06	6.73	8.8	÷.	8.8	7.25	6.B	7.10	117	7.12	98.	2.B
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log P ₁	Ψ	Ý	φ	Ý	Ý	Ý	Ý	9.9	Ý	Ý	Ÿ	Ý	Ý	φ	-7.	Ý	7.	-7.	-7-	Ý	-	Ÿ	-7.	Ÿ	-7.(φ	Ý	-7	φ	-	-	-7.	-7.0	-7.
E E	67.1	48.2	6.0	51.3	57.1	57.6	35	8	8.7	61.3	61.6	3	8.8	8.0	25	65.2	86.1	8.1	8.8	67.0	67.1	67.2	67.2	\$	70.2	5. 0. 1.	1.5	73.2	73.3	73.3	73.3	74.4	74.5	75.0
2	~	•	S	S	S	•	~	•	~	9	-	•	S	•	~	•	•	-	5	S	S	•	•	*	_	S	~	•	~	~	•	**	~	_
•								=																										
•	5.5	7	8.9	75.1	4.17	100	2.67	109.6	33	86.9	5.7	5.40	6.13	1.8	3.5	8.7	2.6	2.5	8 .0	13.33	2.76	12.57	2.8	3.70	3.60	23.	1.27	31.95	4	3.8	14.	10.20	160.0	0.465
P. 6.	*0.76	*1 .1	\$0.34	£0.36	±0.13	£0.73	£0.0	£0.15	£0.3	£0.30	\$0.0	\$0.14	£0.15	19.04	£0.0	£0.10	£0.0¢	\$0.0Z0	÷0.3%	£0.03	£. 3	±0.27	ð.	6.9 6.9	£0.76	, 1, 1,	19.Q	50.00 00.000	<u>۾</u>	± 0.16	±0.052	₽. ₽	\$9.0 \$	11.69
44/4t /esc ²)	-20.71	-17.4	11.92	-10.69	- 9.83	- 6.59	-10.07	- 5.61	- 4.25	- 6.25	- 3.16	- 7.30	- 5.00	-14.14	- 2.291	- 42	- 2. K	. 2.84	37.1	98.7	- s. 8	- 2.84	- S. S.	- 7.08	- 4.54	- 4.86	- 8.02	- 1. ics	- 3.85	- 3.59	- 1.307	- 2.07	- 0.437	- 9.8
, dr (ba/aac) (ba/								X 69								4																		_
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¥r.	1967	190	180	1961	285	190	1541	1967	1939	281	133	190	181	1937	1961	18. 28.	35	1947	190	1961	181	3	3	1937	1981	947	351	1947	261	ž	28	1927	ž	1937
Plate	PA 10557	AI 41417	FA 10557	AI 41417	AI 39337	AI 41417	FA 5202	PA 10557	FA 2602	7585		•			-			¥ 488		-				AI 33580	AI 33553	AI SEO!	AI 39253	A 41394	1. STTT	AI 39265	P. 10557	AI 33:53	AI 34206	1 1 2 2 3
ż	1539	S	1539	1539	1243	1539	1068	53	818	1205	815	154	1068	25	150	1243	120	150	152	15E	7117	106	1112	8	217	198	1241	3	1180	285	153	9	202	169

Data Used in the Compatation and Analysis of

Amorphoric Densities in Massachusette (H < 30 Ks)

Semi	Gentald	Siaid	Tenrid	Teerid	Teerid	Sporadic	Tourid	Tener id	Sporedic	Tourid	Sporadic	Tearid	Sportedic	Lyrid	Sporadic	Sporadic	Tourid	Cardaid	Wirginid	Coming	Sportedic	Feerid	Tear id	Teurid	Sporadic	Sporadic	Sporadic	Perseid	Sporadic	Speradic	yrid	Sporadic	Sporadic
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log Pa	-7.27	8 .9	-7.15	-7.18	-7.24	7.7	-7.37	4.7-	38	-7.47	-7.75	-7.61	-7.67	-1.01	-7.38		-1.46	-7.51	-7.46	-9.15	-7.88	-7.79	S:-	-7.43	-7.73	7.1	÷.14	9 .10	-7.88	-7.78	3.7	7 7	20.
.گر	, O.	ā		8	+.71	8	•.15	a :	3.	8.	<u> </u>	 8	18	3.	*.1¢		÷.19	÷.8	8.	57	£.	<u>-</u> .	•.28	+.35	s 8	•.13	٠. پر	8.		5.	•·18	21	2 3
4	::	ä	¥.	÷	\$	5.	•.31	.73	1.23	•.25	÷.	•.13	2.	•. 51	ij		+.39	•.26	•.17	×.	8	.2	ij	\$	÷.26	•.39	8.	•19	X :	.38		9.5	
log P,	2.8	28	-7.05	-7.06	-7.11	-7.49	-7.29	-7.33	7.0	-7.38	-7.5S	2.7	-7.68	-7.19	-7.33		-7.35	-7.53	2.5	4 .17	-7. 33	-7.68	-4·	-7.35	-7.75	3.7-	4 .18	۲.3	-7.76	 8	-7.76	6. 17	-7.75
4	٠.07	3	£.	3.	ä	5.		•.25	8.	ij	5.	• 14	8.	3.	8.		. X		2.	21	8	÷.x	3	.57	ä		٠.	2.	3.	3.	3.	•.25	2 x
log p.	27.33	-7.01	-7.14	-7.15	-7.19	-7.49	±7.50	-7.X	2.5	-7. 4 5	-7.57	-7.52	2.7-	-7.25	-7.41		¥	-7.8	-7.65	8 7	-7.80	-7.48	₽./-	-7.X	-7.69	-7.61	8.	-7.85	-7.70	57.7-	-7.70	8.8	-7.69 -7.74
Δı	01	×	ij	28	ä	÷.10		÷.	.31		•.24	÷.15		+.25	₹.31		3.	÷.19	•.16	£.	3	•.24	3.	3.	3.	3.	+.15	0	\$	÷.15	¥.	•.19	3.8
I'd But	-7.41	-1.8	-7.21	2.7	-7.19	-7.46	-7.X	7.7	-7.28	-7.36	÷.	-7.51	-7. 8.	-7.45	-7. \$	(& &)	-7.4:	-7.60	3.7	9.10	2.7	3.7-	₹ .~	-7.33	-1.91	 8	-7.98	₽ .15	-7.65	등 주	-7.91	8	-7.7- -7.41
m J	36.0	76.1	16.4	4.0	1.1	78.1	78.7	78.8	3.0	79.0	79.7	79.4	7.6	98	80.2	\$0.4	88.5	81.2	81.4	81.5	1.7	4.4	8.0	3 .4	E .7	E .7	85.5	8	8.2	%	17. S	98.0	8 8
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•	135.8	2.37	1.81	18.5 2	0.23	11.76	<u>ه</u>	2.8	14.3	10.4	15.5	12.4	33.81	1.25	7.37	14.88	1.18	4.0	3.45	8	8.15	0.886	3.8	179.8	1.61	0.462	£3.67	0.07	0.551	5	S. &	o. \$	16.17 0. 45
p.e.	90.04	11:4	\$.¢	£0.36	5 .5	£0, 19	±0.15	5 0.10	£0.03	£0.0g	£0.04	90.0g	£0.880	8	\$ \$	£0.22	8 .8	£0.61	8. 8	£0.2%	6.08 88	£0.16	£0.61	19.0g	27.04	3 7	£0.0%	11.2	£0.26	£0.24	€ €	£0.31	±0.069 ±1.9
44/4t _/amo ²)	8.	7.7-	5	-7.62	₹	-1.19	7.7	-2.65	8	-1.87	6 8	-1.141	8	-7.57	7.7	28 †	-3.E	-7.0	-1.62	-1.53	89. ℃	-2.20	-7. SZ	\$ 주	5.8	-2.51	9.18	9 9	-i %	7.7	-1.82	-1.31	6.0.6 84.0.6
(lm/mc)(36.07	35.61	8.00	33.25	2 .2	28.05	3.6	27.62	20.31	30.4	20.39	2 2 3	27.88	3	20.03	19.75	8	36.35	S Si	8.8	29.21	30.65	31.27	26.2	S. 12	2 6. £	2 2.8	3.S	26.77	S. 58	19.4	¥.73	8 8 8 3
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Plate		PA 10559			_			AI 34319	A STITL	AI 33580	FA 7010			B 238	_	A1 38643	AJ 23564	AI 41418	AI SAS91	AI 33696	AI 39253	AI 22906			_	_				AI 36463			A1 33500 FA 10576
į	1539	1550	<u>8</u>	E	3	126	5	£	1180	ğ	1100	2	3.	1065	125	153	5	1550	8	<u>Ş</u>	3	ğ	<u>8</u>	Š	ĕ	8 6	3	1173	8	161	1065	32	5 3 8

TABLE IV b

Data Used in the Computation and Analysis of

Atmospheric Densities in Massachusetts (H > 85 Km)

Shorer	Perseid Sporadic Sporadic Sporadic Sporadic Sporadic Orionid Perseid Perseid Perseid Sporadic Perseid Sporadic Sporadic Perseid Sporadic Perseid Sporadic Perseid Sporadic Perseid Sporadic Perseid Sporadic Perseid Sporadic	Leonia
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	0.02 0.07 0.07 0.07 0.09	U.843
p.e.	8.01.00.00.00.00.00.00.00.00.00.00.00.00.	\$1.0
dv/dt (læ/sec²)	+ 1.8 - 6.0 - 1.96 - 1.96 - 1.97 - 1.31 - 1.31 - 1.31 - 1.31 - 1.31 - 1.31 - 1.31 - 1.32 - 1.33 - 1.33 - 1.34 - 1.35 - 1.35	7.7 -
v (laz/sec)	8. 25. 25. 25. 25. 25. 25. 25. 25. 25. 25	(1.1)
8	213 22 23 23 23 25 25 25 25 25 25 25 25 25 25 25 25 25	210
F.	416. 82. 82. 82. 82. 82. 82. 82. 82. 82. 82	•10.
Yr.	1942 1942 1943 1944 1944 1945 1946 1946 1946 1946 1946 1946 1946 1946	7230
Plate	A 10243 A 11244 A 11394 A 13612 A 36602 A 36601 A 36601 A 10575 A 36604 A 10576 A 10576 A 10577 A 3660	C7040 TV
į	1469 1544 1173 1005 1005 1005 1005 1006 1009 1009 1009 1009 1009 1009 1009	72.

Data Vood in the Compatition and Analysis of

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matties in Ner-Maxico. ($B < 90$ Ka)	. In p. H. long $ ho_1$ Δ_1 long $ ho_2$ Δ_2 long $ ho_3$ Δ_3 lo	11 1 61 2 56 8 11 A15 - 10 - 428 138	10 10 10 10 10 10 10 10 10 10 10 10 10 1	12 2 57.1 -6.75 -47 -6.71 -48 -6.5224	11 3 57.4 -6.7343 -6.7040 -6.4019	11 3 51 5 51 52 51 11 5 51 11 11 11 11 11 11 11 11 11 1	10 3 56.1 6.33 .00 6.22 6.05 -5.95 6.36	8 2 59.6 -6.49 10 -6.42 03 -6.13 26	13 4 60.4 6.7330 6.6473 6.4603	11 3 60.8 -6.55 11 -6.57 13 -6.32 12	72 6 61.0 6.63 +.02 6.3807 6.88 +.37	12 12.0 - 12.	15 4 61.9 45.6 51.4 45.8 5.17 45.6 5.85	11:- 60:0			8 27 2 37 2 37 37 37 37 37 37 37 37 37 37 37 37 37	25 4 67.1 45.2515 45.2517	24. 5 67.8 -7.3557 -7.3150 -7.2645	21 5 68.3 -6.6123 -6.7212 -6.5130	16 2 68.7 -6.84 •.03 -7.0417 -6.9205	12 2 68.7 -6.8304 -6.8902 -6.7314	50.4 20.5 50.1 50.5 50.5 50.5 50.5 50.5 50.5 5	91 19 3 69.3 -7.0716 -7.1322 -7.0312 -6. 2 21 6 69.7 -7.1118 -7.0411 -6.2201 -6.	12 5 70.0 -7.2125 -7.2933 -7.2479	25 4 70.1 -6.81 +.15 -7.1317 -7.0007	17 3 70.1 -6.6432 -6.9701 -6.8313		N 2,7- 2,- 2,7- 2,7- 2,1- 2,1- 2,1- 2,1- 2,1- 2,1- 2,1- 2,1	10 \$ 72.1 -7.5746 -7.6554 -7.7059	23 4 72.1 -6.9714 -7.2514 -7.1908	16 4 72.1 -6.9417 -7.2312 -7.1605	37 4 73.4 -7.4525 -7.6545 4 77	19 6 74.4 -7.08 •.20 -7.3406 -7.3002	30 6 74.4 -7.16 -12 -7.42 -34 -7.4012	13 2 74.6 -6.8049 -7.0029 -6.8742	11 3 74.6 -7.5 - 127.1 - 1.5 - 1.9 - 1.	01.0 02.7- 20.0 62.7- 20.0 22.7- 7.27 9 52.		12 Z 75.2 -6.81 +.54 -6.94 +.34 -6.87 - 1.47
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for Mari	= Ĵ	4 13	8	2	15	6.72	3	39.6	8	8.	2.0	91.0	6.5	;	3	3 3	3	67.1	67.8	3.3	3	25.7	5	S S	2	10.1	25	7.5	1.9	1.1	1.5	1.	7.5	4.	7.7	4.6	4.6	7		(3.Z
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4	•	2 60	0	9	1.2	-	21.13	2.5	3 7.04	8.33	S	\$7.74 2.24	200	2	2 9	2 9	4	8		£0.22	\$0.65	£0.76	17.5	5 6 3 8	60.19	£0.1¢	20.35	2 9	9	67.04	Š Ž	£0.19	20.2	10.0	£0.06	8		를 : 얼 :	2 S	2
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Atmospheric Danelties in New Mexico (# < 30 Es)

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10g P ₁	*	7.7	-7.16	7.7	7.	-7.3	3	2.7	-1.0	7.13	7.4	-1.8°	3.7	9.7	3.	9	9	 8	¥.~	3.7	\$	8.7	-7.41	4.15	3.7-	-7.19	8.7	- 7.€	3.5	3.7	S .	*	-7.93	 S	9 7	8 :~	7.7			۲.3		 	18
≖Ĵ	×	×	7	75.6	7	2	7.9	7	3.5	₹.S	3.6	3	۲. د.	F.	-		2 .5	78.5	3.6	39.8	£.9	8	2.5	2.5	2.5	79.3	6.6	8	80.7	8	9.0	0.5	3.	9.18	61.7	:	<u>.</u>	2.2	2.2	2.2	2.3	2	7.0
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	17.11	-14.0	2	-13.4	- 1.17	- 2.2	5.5	-10.08	- 6.57	- 2.13	- 2.2	2.2	3.0	=	20.	- 0.71	- 2.15	- 1.67	- 8:	-1. .	85 -	87	-11.6	- 0.66	- 2.70	- 1.765	- 1.38	- 1.470	3.°	. 3.39	.3.8	1.39	2.2	- 0.72	- 0.465	- 2.17	- 6.0	7. 0 -	- 0.33	- 3.4	4.0.	8	- 0.12
(jm ,	2	3	8	52.73	8	25.39	8	3	3 .8	23.23	z Z	Z.X	17.8	8	2.2	B	N N	8	26.18	K.	31.10	21.12	30.00	K.3	21.79	19.33	19.08	19.61	3. 8.	8	S S	2	= : ::	31.16	3.5	X.55	8.9	35.48	Z.	39.38	19.83	20.19	35.85
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≖Ĵ	7.0	8		20.0	2	2.2	8.3	2.5	2.7	8.9	ĭ	Ę	ž	ž	ž.	3	3	3	S,	15.3	15.6	25.6	5.6	5	3	3	8.5	%	1.5	1.1	17. S	1.7	2.7					9				1	8	3.6	3.0	?
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•	2	5	3	3	9	ÿ.	2	9.X	0.17 17	97.0	E.	9.11.	1.1	S S	9.1	O.1.5	1.17		1.1	14.80	51.3			0.217		5.43		-: *:	22.2	9.45	0.457	9	Z,	3			8	0	777	0.65	0.X	2	3.	6	6	3,
į	77.0	6	99.99	8.7	8 .7.	5 0.10	8.10	8	1.13	8	8.9	4.2	10.13	8.13	2.2	3.5	8.8	±0.0013	8.8	8	A. 8	4	7	1.16	9	8:18	8.0	\$0.015	.S.	\$0:04	8 .5	\$ 6.0	2.5	8	3 2			đ	1	7	7	Z,	1.8	#0.X	7	7.7
() ()	2.5	3	- 0.0	- 1.73	. 3.8	男.0.	- 0.55	1.1	X -	- 1.17	- 1.915	- 2.9	. 1.8	- 1.51	- 2.6	- 2.4	- 0.613	- 6.156	- 6.75	- 0.167	- 0.861	- 3.14	- 3.13	- 2.5	- 0.0	- 3.6	87.0	- 0. XS	9	- 1.X	- 1.8	- 1.18	7	- 6.119					0.0076	- 6.19	- 0.355	1.1.	- 0.93	- 1.8	- 2.3	- 0.X
. j	25.66	2.2	3.5	スコ	n.c	X. IS	19. K	Z	KX	Z, Z	8	3	X.13	27.76	8	3.3	21.76	Z	X	20.5	M.K	8	2.5	7	7	8.14	27.10	23.88	20.74	31.6	X.*	4	X.X	# F	į	3		3	7	19.61	39.65	4	3.6	65.17	8.0	21.73
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Ţr.	88	8	30	388	8	Ĭ	991	3	Ĭ	38	8	3	8	1960	35	861	98	3	8	8	3	9	35	3	2	98	8	350	3	88	367	3	3					98	3	186	3	9561	1980	3	8	80
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ė	22	70	N N	E	1551		S N	E 361	255	186		N X	FEE	198	2007	2557	2	X10	K SX	70 101	1680 A.	269	Z CZ	7127	N SUN	203	N SN	žž.	178	1947	1937 EA	171				200	128	1910	Z S	1910	2011	7 62	2	1760		1551

TABLE V b

Data Used in the Computation and Analysis of Atmospheric Densities in New Mexico (H > 85 Km)

Shorer	Tearid Sporadic Sporadic Caninid Caninid Caninid Sporadic Tearid Sporadic	Sporadic
ρ ₂ (×10 ⁹)	1. 25 2. 25	ઝ
ρ ₁ (×10°)	13. 23. 23. 23. 23. 23. 23. 23. 23. 23. 2	ង
E ∃	88 88 88 88 88 88 88 88 88 88 88 88 88	7.16
•	2.0.1.2.0.0.2.0.0.2.2.2.2.2.2.2.2.2.2.2.	15.5
•	** * * * * * * * * * * * * * * * * * *	2
•	13.4 14.1 15.0 16.0 16.0 17.0	0.017
P. e.	6.03 6.03	€0.70
dv/dt (lan/sec#)		- 0.38
(ba/sec)	%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%	64.68
•	28	19
ř	និងខ្មែននិងខ្មែននិងជាជានិងខ្មែននិងនិងខ្មែននិងនិងនិងនិងនិងនិងនិង	8
Yr.	25	1950
Plate	N 1262 N 1262 N 1260 N 1267 N 1267 N 1267 N 1267 N 1215 N 1217 N 1217	A 1525
No.	2620 K 2620 K 2620 K 2620 K 2620 K 2620 K 2610 K 26	1918 K

TABLE V b

Data Used in the Computation and Analysis of

Atmospheric Densities in New Mexico (H > 25 Km)

ė	Plate	Ĭr.	Ė	•	(lm/eec)	dv/dt (bs/esc ²)	••	=		•	æ <u>(</u>	ρ ₁ (×10°)	$(\times 10^9)$	Shower
88	1723	1950	Š	110	37.53	- 1.53	40.40	0.123	12	8.8	91.8	3.2	5.2	Lyrid
2548	9	280	3	3	8.8	- 0.323	.0.024	0.73	\$	8.3	91.9	2.2	5.6	Genius
1740	KA 1319	1949	86.	287	65.45	- 1.12	£0.29	0.420	19	15.2	2.1	2.0	7.7	Sporadic
22.39		155	5	S	4.17	- 1.17	\$2.5%	1.70	61	1.6	2.3	7.1	10.5	Sporadic
202	PA 12162	198 08	9 19.	ជ	8 .8	- 2.40	12.03	0.584	Ħ	11.4	3 .4	5.5	11.1	Perseid
1728		186	8	8	2.14	- 2.66	11.67	0.62	61	0.0	8.5	\$.	ಸ	Sporadic
222		28 28 28	3.	7	51.21	- 3.1	11.4	0.89	=	0.1	23.7	2	17.	Sporadic
2024		<u>8</u>	8	219	89.99 99.99	- 6.07	*1.15	0.0128	Ħ	8.6	ż	9.0	4.9	Persoid
9998	-	1981	ş.	~	2.3	- 1.35	±0.96	0.063	61	1.0	98.3	3.1	4.5	Sporadic
1910		5 5 5	8	<u>\$</u>	\$0. 4 0	- 2.34	\$0.0	0.114	X	S.	95.4	4.5	7.6	Lyrid
1991		1950	ğ	8	30.05	- 2.19	£0.59	0.0433	11	5.7	9 8. 4	3.0	5.1	Lyrid
2313		25 25 26	3.	3	4. %	- 0.81	€0.79	0.11	61	1.0	8.0	2.0	3.0	Sporadic
2024		55 55 55	8	219	61.11	- 3.34	#1.40	0.0185	3	3.5	8.7	2.3	4.7	Perseid
2313	-	28	3.	3	4. 68	+ 0.23	*1.47	0.13	25	0.3	%	- 0.6	- 0.9	Sporadic
1910		1950	ğ	8	50.57	- 2.33	\$0.04	0.125	8	\$	%	4.6	7.8	Lyrid
2002		1980	<u>§</u>	218	57.57	- 5.5	13.4	0.0364	ដ	0.3	9.0	5.5	10.6	Perseid
1755		1949	5 9.	H	83.83	• 4.0	13.6	0.046	2	0.3	97.2	- 3.5	- 7.4	Sporadic
222	-	1980 1980	3.	3	20.69	- 1.37	₹0.54	16.0	ដ	5.5	97.3	5.1	8.6	Sporadic
1698		186	E.	316	8.8	- 0.41	£0.65	0.054	11	9.5	97.5	0.39	0. 8	Sporadic
9 0 02		1950	919 .	ឆ	60.70	- 0.89	£0.69	0.107	11	4.6	77.5	1:1	2.2	Sporadic
1991	ICB 1575	386	28	8	S. 53	- 2.30	±0.35	0.0537	×	17.9	7.16	3.4	5.8	Lyrid
2061	FA 12064	55 55 55	\$	E	3 .8	- 9.9	£9.7	0.024	13	0.1	8 .6	7.7	16.	Sporadic
1898	TS11 80	195	E.	316	2.8	• 1.7	11.4	0.086	91	1.6	8.3	- 1.8	- 3.8	Sporadic
1755	IV 1171	1949	5	ដ	2.9	. 7.0	*1.5	0.118	ជ	6.0	₹	8.4	9	Sporadic
205	AI 42782	9 <u>26</u>	8	218	27.86	- 2.15	£0.66	0.0450	1	9.9	8 .8	2.3	4.4	Perseid
30 61	AI ethi	2 <u>8</u>	9	173	61.03	- 7.7	*4.0	0.029	13	0.3	8.0	9.	13.0	Sporadic
ZZ ZZ	FA 12423	1950	ä	3	59. 17	+ 3.89	±0.45	11.9	X	•	101.1	(- 2.5)•	- (6.4 -)	Sporadic
2679	69K	56 56 57	ş	ğ	8.8	- 0.21	£0.33	0.074	Ç	22.5	102.5	0. M	°.	Sporadic
2679	3	1980 200	3	ğ	52.59	- 0.72	£0.55	0.081	19	6.9	182.8	1.12	 %	Sporadic
2181	KA 2184	198 98	2 .	317	7.5	- 4:1	42.6	0.041	11	6.0	102.9	2.8	6.7	bino
2179	0867 IV	35	Ë	316	71.20	+ 0.68	* 1.52	0.048	ជ	2.5	183.3	- 0.49	- 1.16	Leanid
223	AI 43021	2 <u>8</u>	ă	Š	SS. 33	• 0.693	±0.067	12.3	#	•	104.4	[- 4.6]•	- 0.0]	Sporadic
2176	AI 42987	<u>8</u>	E.	316	70.21	• 0.23	₹0.80	0.054	೩	10.3	107.2	- 0.18	- 0.42	Sporadic
2468	AI 42916	5 5 8	.7 80	ž	65.11	- 0.516	±0.278	0.203	×	X .2	108.6	٠. د.	7.5	Sporadic
2468	FA 12314	1950	.789	ž	65.33	- 0.65	£0.70	0.0	22	4.8	109.0	0.8	2.0	Sporadic

. See remarks about this meteor in the text, section 1 .

small relative error when the point we consider is near the beginning or the center of the trail, but may become quite important when we have to compute a mass near the end of the meteor's visible path. It is, therefore, necessary to establish a weight function which will take into account all these factors.

If the scatter in the computed atmospheric densities were due only to observational errors, a weight function could be established rather simply on the basis of the probable errors of du/dt and m. We must not forget however, that even densities computed from exact data (i.e., with probable errors equal to zero) would presumably show a scatter due to different shapes and densities of the individual meteoroids and to unaccounted day-to-day variations in the upper-atmospheric layers. While a weight function based entirely on probable errors would be justified for data with very low accuracy, there is the danger that it would lead to dangerous relative overweighting of more accurate data. As we see, on ideal weight function must follow the probable-error rule in the low-accuracy domain and reach an upper (saturation) value when the probable errors decrease beyond a certain limit. The weight functions which were finally adopted are certainly far from being ideal and would be difficult to justify on a rigid anslytical basis; they do, however, satisfactorily accomplish their assigned task, so we feel they can be presented without further apologies.

Whenever values of $\log_{10} \rho$ were analyzed (i.e., for heights lower than 85 km), their weight ρ was assumed to be

$$p = 10 \ \psi \ (\frac{\dot{v}}{\epsilon}) \ \dot{\psi} \ (n-3) \ (\frac{\pi}{\pi_{\infty}})^{\frac{1}{2}} \ ;$$

$$\psi = \frac{1}{2} \left[1 + \operatorname{erf} \left(2 \log_{10} x - 0.7 \right) \right] \ . \tag{3}$$

Here \dot{v}/ϵ is the ratio of the acceleration to its probable error, n is the number of ahutter breaks used in the least-squares solution (which contains three unknowns) and n_m the mass of the meteor before it entered the atmosphere. In the expression $r \psi$, x may be either \dot{v}/ϵ or n-3. The weights thus computed range from 1 to 9 and virtual all densities with $\dot{v}/\epsilon < 2$ have weight zero.

For heights above 85 km, where weighted means were taken in ρ rather than $\log_{10} \rho$, the weight function was

$$w = \left[\frac{\pi}{4}(n-3)\right]^{\frac{1}{2}} \operatorname{erf}\left(10^{-18} \epsilon_{\rho}^{-2}\right),$$
 (4)

where $\epsilon_{\rho} = (\rho/\dot{\nu})\epsilon$. The meteor mass does not appear in this formula because at great heights all accelerations are computed from centers or early parts of trajectories.

The separation of the observational material into two sections, below and above 85 km, appeared to be necessary in view of the great disparity in the accuracy of the data. Below 85 km the observed decelerations are generally much larger than their probable errors and there are virtually no negative densities. Above 85 km it becomes increasingly difficult to get reliable decelerations and negative densities become more and more frequent.

In view of the nearly linear relation between H and $\log \rho$, it appears logical to prefer taking weighted means of these two quantities whenever possible, and this was done for H < 85 km. For H > 85 km advantage was taken of the fact that atmospheric densities in the range 0 < H < 100 km can be represented, within a factor of 2, by the approximation $\rho = \rho_0 e^{-0.138H}$. Weighted means were therefore taken in ρ and in the quantity $e^{-0.138H}$, from which the corresponding mean value of H was computed.

3. Analysis of the New Mexico Data:

The mean atmospheric density profile over New Mexico was assumed to be sufficiently well established on the basis of the results of several high-altitude rockets launched in recent years. N.R.L. dats from four rocket flights were reduced by the writer in 1948⁽¹⁾ and the mean profile derived from them (profile C, Table VI) has been used for the past 4 years at M.I.T. as

a guide in problems involving the upper atmosphere $^{(6)}$. Our chief goal has been to establish an empirical function of the densities ρ_1 computed by equation (1) from New Mexico meteors, which would reduce the computed densities as closely as possible to those of profile C. This same function would then be applied to the densities computed from Massachusetts meteors; the two sets of data would then presumably represent true atmospheric densities and could be analyzed and compared.

A mean density profile for the upper atmosphere over New Mexico has been recently computed by the Rocket Panel* using data from 16 rocket flights⁽⁵⁾. This profile is given in Table VI under the designation of profile R. Profiles R and C are almost identical (they intersect three times in the range from 50 to 100 km) so no appreciable difference in our results could be expected if one were substituted for the other in the analysis.

As a preliminary step, the individual values of $\log \rho_1$ were compared with three different density profiles, given in Table VI. Only points lying between the heights of 65 and 85 km were used in the analysis. This limitation stems from various considerations:

TABLE VI

Atmospheric Density Profiles Used in the Analysis of New Mexico and Massachusetts Meteors (A,B,C) and Most Recent Profile from V-2 Rockets (R). Log₁₀ \(\rho \) in g/cm³.

H (km)	A	В	С	R	ä
45	-5.744		-5.725	-5.667	
50	-6.006	-5.99	-5.971	-5.936	-5.91
55	-6.227	-6.21	-6.193	-6.199	-6.15
60	-6.421	-6.42	-6.409	-6.457	-6.41
65	-6.607	-6.63	-6.647	-6.723	-6.70
70	-6.801	-6.88	-6.955	-7.012	-6.97
75	-7.018	-7.20	-7.324	-7.334	-7.25
80	-7.287	-7.56	-7.700	-7.676	-7.52
85	-7.611	-7.92	-8.062	-8.034	-7.80
90	-7.951	-8.29	-8.400	-8.389	-8.08
95	-8.273	-8.67	-8.717	-8.734	-8.36
100	-8.569	9.04	-9.021	-9.063	-8.66
105	-8.852		-9.314	-9.379	
110	-9.131		-9.598	-9.684	

- A = Profile from Massachusetts meteors, 1949(3).
- B = Preliminary profile from New Mexico meteors, using uncorrected data and $\log K = 0.12$.
- C = Profile from New Mexico rocket flights, 1948(1).
- R = Most recent profile from V-2 rockets, 1952⁽⁵⁾.
- M = Final density profile from Massachusetts meteors, from results of the present paper.

¹⁾ In the height range between 65 and 85 km there is a generous overlapping of meteor velocities and masses. At heights lower than 65 km densities are derived mostly from low-velocity, massive meteors near the end of their trajectories. At heights greater than 85 km high-velocity, low-mass meteors predominate.

The author wishes to express his sincere gretitude to the Rocket Fenel for the permission to use these deta before their publication.

- 2) At heights greater than 85 km individual meteor decelerations become progressively less reliable.
- 3) The introduction of the height in the form of a linear term in the equation of condition precludes taking any wider range in heighta, or the term would lose much of the rightaleance.

Since the exact form of the drag equation (1) is still open to discussion, both the velocity ν and the mass a were inserted in the equation of condition, although it was realized that they are not independent parameters; also introduced were the height H, to account for an error in the slope of the comparison curve, and the fraction of the tropic year to account for a possible seasonal effect. This last quantity was introduced through an auxiliary angle $\phi = 2\pi t/T$ where T is the duration of the tropic year and t the time elapsed since January 1.0 of the year when the meteor appeared.

The equation of condition used had the form

$$\log \rho_1 - \log \rho' = a + b \sin \phi + c \cos \phi + d (\log_{10} \nu - 6) + eH + f \log_{10} m$$

$$= a + M \sin (\phi + a) + d (\log_{10} \nu - 6) + eH + f \log_{10} m.$$
(5)

Here ρ' is the density at the height H in the comparison profile (A, B, or C).

In view of the peculiar distribution of the meteor material, none of the parameters is, strictly speaking, independent of the others: velocities are closely connected with heights, masses with velocities and there is a tendency for low-velocity meteors to abound in the winter season and high-velocity meteors to prefer the summer season. The situation, as will be seen, is particularly bad in this respect for the Massachusetta material. For New Mexico, however, the distribution is much more uniform with respect to all parameters and some sense can be made out of the various correlations. A comparison of the solutions made using the complete equation of condition (4) with others in which the coefficients of log s or of H or both have been put equal to zero is particularly instructive.

A special effort was made by the New Mexico observers to cover the Geminids of 1950, and this resulted in an abnormally high percentage of such meteors in the analyzed material. Since a cursory inspection had shown that the atmospheric densities derived from Geminids seemed to be systematically lower than the average, separate least-squares solutions were made, with and without Geminids.

The results of the various least-aquares aclutions for the New Mexico data are given in Table VII. An inspection of Table VII leads to the following preliminary results:

- a) No matter what comparison profile is taken, and whether or not Geminids are included, the value of d, the coefficient of $\log \nu$, is always -1.1 or -1.2, and exceeds its probable error by a factor of 10. The 1948 analysis of Massachusetts meteora (1) had given for the coefficient of $\log \nu$ the values of -0.91 and -0.79, according to the equation of condition used.
- b) The coefficient of $\log n$, as was to be expected, is of opposite sign with respect to that of $\log \nu$. It is important to note that its magnitude is of the same order as that of its probable error and that, when it is put equal to zero in the equation of condition, the remaining coefficients change but very little.
- c) The seasonal effect is small, with an amplitude which for most solutions amounts to less than one-half of that derived for Massachusetts in a previous analysis [1]. The amplitude is only two to four times larger than its probable error and the phase of the fluctuation is quite different from the one found for Massachusetts. The seasonal effect is, however, atrongly dependent on the comparison profile, and any more definite conclusion on its reality should be deferred until later.

A plot of $\log \rho_1$ against H (fig. 1) shows an excellent agreement with the rocket profile C. On the other hand the resulta listed under A) show that the residuals from this profile are strongly dependent on velocity, and that the cause of this dependence must be sought not in the comparison profile, but in the fundamental equation (i). It appears that the correct value of d is very close to -1.0, which would correspond to an error of one power in v in equation (i). As a second step, then, we tried to make the computed densities independent of velocity by

TABLE VII $\label{eq:analysis} \textit{Analysis of Log} \ \rho_1 \ \textit{for New Mexico Meteors}$

65 < H < 85 Km

No.	Comparision Profile and Data Used	•	ь	С	đ	•	f	M	α	Min	•
1	A, All Meteors	+0.197	+0.077	-0.006	-1.19	+0.0002	+0.032	+0.078	355°	Oct.	7
	•	±0,200	±0.028	±0.020	20.12	:0.0026	₽.019	±0.028	±15	±	16
2	A, All Meteors	+0.184	+0.054	-0.004	-1.22	+0.0004		+0.055	356	Oct.	7
		±0.205	±0.025	±0.021	±0.12	±0.0027		±0.025	±22	*	23
3	A, No Geminida	+0.436	+0.091	+0.017	-1.12	-0.0033	+0.049	+0.093	11	Sept.	21
	.,	±0.231	±0.038	±0.021	±0.12	±0.0031	±0.021	±0.045	±13	±	14
4	A. No Gezinide	+0.333	+0.055	+0.015	-1.18	-0.0016		+0.057	16	Sept.	16
		±0.229	±0.024	±0.021	±0.11	±0.0030		±0.023	±22	*	22
5	B. All Meteors	+0.490	+0.088	-0.032	-1.09			+0,094	340	Oct.	22
	•	±0.053	±0.024	±0.021	±0.12			±0.023	±13	*	13
6	B, All Meteors	+0.013	+0.104	-0.129				•0.165	309	Nov.	23
		±0.018	±0.029	±0.022				±0.025	± 9	±	9
7	B, No Geminide	+0.521	+0.083	+0.017	-1.13			+0.065	11	Sept.	20
		±0.051	±0.022	±0.022	±0.17			±0.022	±15	±	15
8	B, No Geminida	+0.035	+0.094	-0.093				+0.132	315	Nov.	16
		±0.018	±0.029	±0.024				±0.029	±11	±	12
9	C, All Meteors	-1.133	+0.076	-0.009	-1.22	+0.0217	+0.028	+0.077	354	Oct.	9
		±0.207	±0.029	±0.021	±0.13	±0.0027	±0.020	±0.029	±16	*	16
10	C, All Meteors	-1.143	+0.056	-0.007	-1.25	+0.0221		+0.057	353	Oct.	9
		±0.207	±0.025	±0.021	±0.12	±0.0027		±0.025	±22	*	22
11	C, No Gaminida	-0.883	+0.089	+0.013	-1.15	+0.0183	+0.041	+0.089	9	Sept.	
		±0.236	±0.029	±0.022	±0.12	±0.0031	±0.022	±0.029	±14	±	14
12	C, No Gaminida	-0.969	+0.058	+0.012	-1.21	+0.0197		+0.059	11	Sept.	
		±0.233	±0.024	±0.022	±0.12	±0.0030		±0.024	±21		21

This table gives the results of the various least-squares solutions computed, using the equation of condition

-1.05

±0.19

+0.020

±0.024

10 ±11

+0.118 ±0.025

+0.505

±0.058

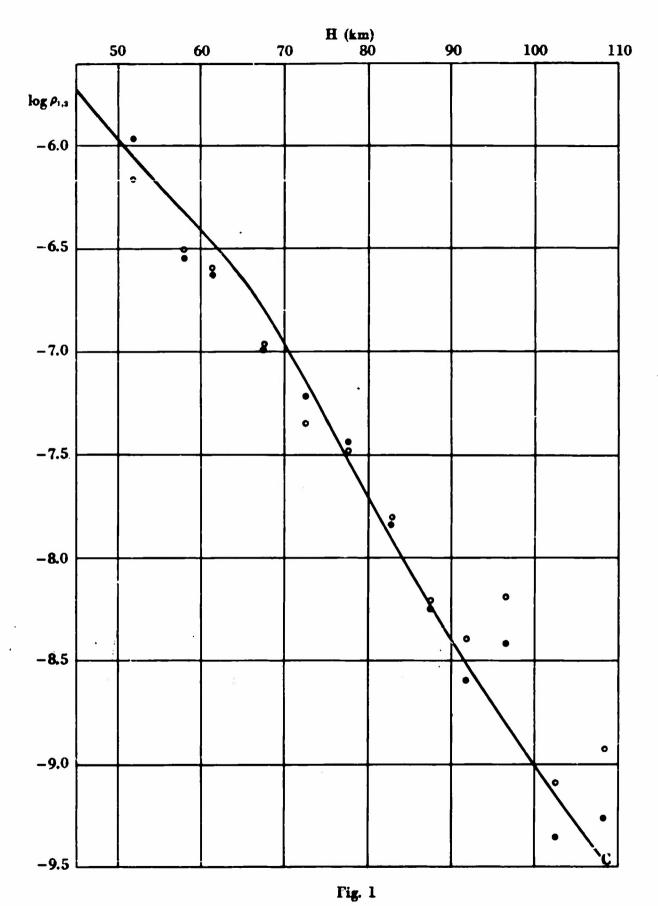
+0.116

±0.025

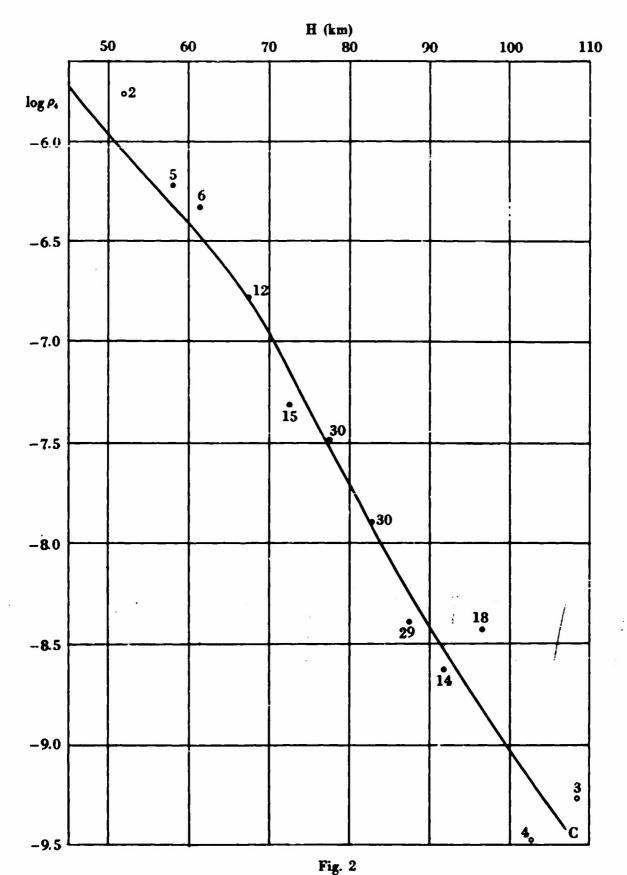
13 C, No Geminida

 $\log \rho_{i} - \log \rho' = s + b \sin \phi + c \cos \phi + d(\log_{10} v - 6) + sH + f \log_{10} m$ $= s + M \sin(\phi + \alpha) + d(\log_{10} v - 6) + sH + f \log_{10} m.$

Here ρ' is the density at the height H in the comparision profile specified for each solution; $\phi = \frac{2\pi}{T}\epsilon$ where T is the duration of the tropic year and ϵ the time elapsed since January 1.0; v is the meteor velocity in cm/sec. and v the meteor mass in grams.



Means of $\log \rho_1$ (dots) and $\log \rho_2$ (open circles) from New Mexico Meteors. The rocket profile C is shown for comparison.



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Corrected Mean Atmospheric Densities $\rho_{\rm q}$ from New Mexico Meteors. The figures which accompany the plotted points indicate the number of observations used to compute the mean. Means from fewer than five observations are shown as open circles. The rocket profile C is shown for comparison.

reducing them, assuming d=-1, to a standard velocity of 30 km/sec. These corrected densities were designated ρ_2 ; we thus have the following equation for ρ_2 :

$$\rho_2 = \frac{v}{3 \times 10^6} \rho_1 \,. \tag{6}$$

In view of the small change in velocity during the course of a meteor trajectory, equation (2) can be written, approximately, as

$$\mathbf{a} = \frac{2}{\tau_0 v^3} \int_{t}^{t} I \, dt = \frac{2}{\tau_0} L \, v^{-3} \,. \tag{7}$$

This means that, for all practical purposes, equation (1) becomes

$$\rho = -\frac{2K}{\tau_0} L^{\frac{1}{3}} v^{-4} \frac{dv}{dt}. \tag{8}$$

The change from ρ_1 to ρ_2 would imply a change from ν^{-4} to ν^{-3} in this equation. Were it not for the results listed under b), we could be entirely agnostic and ascribe the blame for this change indifferently to an error in the power of ν in (1) or in (7). Ascribing the blame to equation (7) would mean making m independent of velocity. Then all meteors having the same brightness would have the same mass, and the thermal theory of meteor light would have to be abandoned.

There are some indications that this is not the case. The excellent agreement of theoretical light curves with observations (2) and the behavior of low-velocity meteors (2) would indicate a definite dependence of masses on velocity. Apart from this, the results listed under b) show that there is at best only a doubtful dependence of density residuals on the mass as computed by (2). It looks as though most of the trouble arises from the form of the drag equation.

Means of $\log \rho_2$ against H are shown in Table IX and in Fig. 1. Since greater velocities are observed at greater height, the velocity correction ($\log \rho_2 - \log \rho_1 = \log \nu - 6.477$) increases with height and, conforming to expectation, the slope of $\log \rho_2$ becomes smaller than that of $\log \rho_1$, which agreed with the alope of the rocket profile. We are thus left with a set of "densities" which are practically independent of velocity, as we desired, but do not agree with the known density profile. It would seem that the introduction of a scale factor R somewhat larger than unity would restore the proper slope. Densities ρ_3 computed by the equation

$$\log \rho_3 = A + R \log \rho_2 \quad (A = \text{constant}) , \tag{9}$$

would leave residuals from profile C which would presumably prove independent of velocity and height, but would still be affected by a possible seasonal fluctuation. The final densities $\rho_{\rm q}$, corrected for seasonal effect, were computed by least-squares, using the equation of condition

$$\log \rho_{\rm c} = A + R \log \rho_2 - B \sin \phi - C \cos \phi$$

$$= A + R \log \rho_2 - M \sin (\phi + \alpha)$$
(10)

where $\rho_{\rm c}$ is the density on profile C at the height corresponding to $\rho_{\rm c}$. The values of $\rho_{\rm c}$ computed by (10), once A, B, C and R have been determined. The presence of a scale factor $R \neq 1$ must be considered as a purely empirical device to make the computed densities agree with the observed rocket profile. A theoretical explanation is not attempted at this stage.

Since a large range in heights is necessary for an accurate determination of the scale

factor R, the least-squares solution was extended to all meteors with heights lower than 90 km, Geminids excluded. The results are given below:

A = +2.059 ± .288 B = +0.056 ± :028 C = -0.034 ± .024 R = +1.275 ± .151 M = +0.066 ± .027 a = 329° ± 22° (Minimum density: Nov. 3 ± 22^d)

Within the same range of heights as that of the least-squares solution there are 25 densities determined from Geminids. When $\log \rho_{\rm q}$ is computed for them, they yield residuals $\Delta_{\rm q}$ from profile C, whose weighted mean is -0.10. We consequently assumed the value of +0.10 to represent the mean correction to $\log \rho_{\rm q}$ as determined from Geminids. Mean values of $\log \rho_{\rm q}$ in function of height are shown in Table IX and Fig. 2. For heights above 85 km, where means were taken in ρ rather in $\log \rho$, the mean of the individual seasons! corrections was applied to each mean point. For the same points the correction to $\log \rho_{\rm q}$ for the presence of Geminids was taken as +0.10 times the ratio of the sum of the Geminid weights over the total sum of the weights within the 5 km range of heights considered.

As a final check on the non-dependence of ρ_{ij} from meteor velocities, the coefficient of correlation between $\log \rho_{ij}$ and $\log \nu$ was computed for all meteors with H < 90 km (corrected Geminids included) and found to be equal to 0.002.

The amplitude of the seasonal fluctuation which results from this least-squares solution is only twice as large as its probable error. In view of the possibility that the seasonal effect might be a function of height and that the picture might become somewhat blurred if the analysis were made over too large a range in height, another solution was computed for heights between 65 and 85 km, Geminids excluded, using a slightly different method. Log ρ_3 was computed by equation (9), using A = +2.05°s and R = 1.27°s, and the residuals Δ_3 from profile C0 were analyzed for seasonal effect. The result of the least-squares solution is given below:

$$\Delta_3 = -0.0096 + 0.083 \sin (\phi + 359°.5)$$
 $\pm .0083 \pm 0.040$
(11)

As we see, the results are not too dissimilar from those obtained using all meteors with H < 90 km. The amplitude is 25% larger, but still only twice as large as its probable error, and the minimum density falls on October 2, or one month earlier. In our judgment it would be a little premature to give a definite pronouncement concerning the reality of the effect.

4. Analysis of the Massachusetts Data:

Values of $\log \rho_1$ for most of the Massachusetts meteors have been published in Technical Report No. $2^{(1)}$. For a number of meteors the data are presented here in a somewhat different form. Previously, when decelerations were determined on two plates for the same meteor, their weighted mean was used in most cases to compute atmospheric densities. To make the Massachuaetts material entirely consistent with that from New Mexico, decelerations from different plates are here treated aeparately.

For a comparison with the New Mexico data we must compute ρ_2 and ρ_3 according to equations (6) and (9), analyse ρ_3 for seasonal variation and finally obtain a profile of ρ_4 . Before we proceed to do this, however, it will be instructive to look at the results of some preliminary analyses of ρ_1 which were undertaken, a little haphasardly, in the early stages of this investigation. Table VIII summarises the results of least-squares solutions computed exactly in the same manner as for New Mexico meteors, with the equation of condition (5), using two different comparison profiles. This table should be compared with Table VII.

TABLE VIII $\textit{Analysis of Log} \ \rho_1 \ \textit{for Massachusetts Meteors}$

65 < H < 85 Km

No.	Comparison Profile and Data Used		ь	с	d	•	f	M .	α	Min	
1	A, All Meteors	-0.734	-0.104	-0.041	-0.762	+0.0109	-0.002	+0.112	202°	March	11
		±0.249	±0.036	±0.037	±0.212	±0.0035	±0.078	±0.036	±19	±	19
2	A, All Meteors	-0.736	-0.104	-0.042	-0.759	+0.0109		+0.112	202	March	1
		±0.244	±0.034	±0.036	±0.204	±0.0034		±0.035	±19	±	1
3	A, No Geminide	-0.872	-0.091	+0.001	-0.593	+0.0122	-0.045	+0.091	179	April	
		±0.246	±0.034	±0.039	±0.281	±0.0037	±0.031	±0.034	±25	±	2
4	A, No Geminida	-0.904	-0.108	+0.006	-0.412	+0.0109		+0.108	177	April	(
	·	±0.244	±0.032	±0.039	±0.251	±0.0036		±0.032	±21	±	
5	A, No Geminida	-0.246	-0.159	-0.014	-0.050			+0.160	185	March	2
		±0.114	±0.030	±0.040	±0.228			±0.030	±14	±	1
6	C, All Meteors	-2.220	-0.111	-0.042	-0.775	+0.0344	+0.009	+0.119	201	March	Ľ
	•	±0.255	±0.036	±0.038	±0,217	±0.0035	±0.060	±0.037	±18	±	1
7	C, All Meteors	-2.210	-0.108	-0.041	-0.791	+0.0345		+0.116	201	March	Ľ
		±0.250	±0.035	±0.037	±0.209	±0.0035		±0.036	±18	±	19
8	C, All Meteors	+0.066	-0.266	-0.156	-0.070			+0.30s	210	March	:
		±0.143	±0.045	±0.051	±0.281			±0.047	± 9	±	!
9	C, No Geminide	-0.286	-0.275	-0.052	+0.661			+0.280	191	March	22
		±0.165	±0.043	±0.056	±0.330			±0.043	±11	±	13

For detailed explanations, see bottom of Table VII .

TABLE IX

Mean Atmospheric Densities from New Mexico Meteors

(Weighted Means within 5-Km Groups; Geminid Densities Corrected)

H (km)	$\log \rho_1$	Δ,	log ρ_2	Δ,	log ρ ₃	, ^Δ ,	log ρ_{ij}	٨,	n
51.9	-5.97	+0.09	-6.17	-0.11	-5.81	+0.25	-5.76	+0.30	2
58.0	-6.55	-0.23	-6.51	-0.19	-6.26	+0.06	-6.22	+0.10	5
61.3	-6.63	-0.16	-6.60	-0.13	-6.38	+0.09	-6.33	+0.14	6
67.4	-6.99	-0.20	-6.97	-0.18	-6.84	-0.05	-6.78	+0.01	12
72.6	-7.22	-0.07	-7.35	-0.20	-7.32	-0.17	-7.31	-0.16	15
77.6	-7.44	+0.08	-7.48	+0.04	-7,49	+0.03	-7.48	+0.04	30
82.7	-7.84	+0.06	-7.81	•0.09	-7.90	0.00	-7.69	+0.01	30
87.5	-8.25	-0.01	-8.21	+0.03	-8.41	-0.17	-8.39	-0.15	29
91.7	-8.60	-0.09	-8.40	+0.11	-8.66	-0.15	-8.63	-0.12	14
96.5	-8.42	+0.39	-8.19	+0.62	-8.39	+0.42	-8.43	+0.38	18
102.6	-9.36	-0.18	-9.10	+0.08	-9.55	-0.37	-9.48	-0.30	4
108.3	-9.27	+0.23	-8.93	+0.57	-9.33	+0.17	-9.27	+0.23	3

 Δ_1 , Δ_2 , Δ_3 , Δ_k are residuals of the respective $\log \rho'$ s from the Rocket Profile C (Table VI)

Mean Atmospheric Densities from Massechusetts Meteors

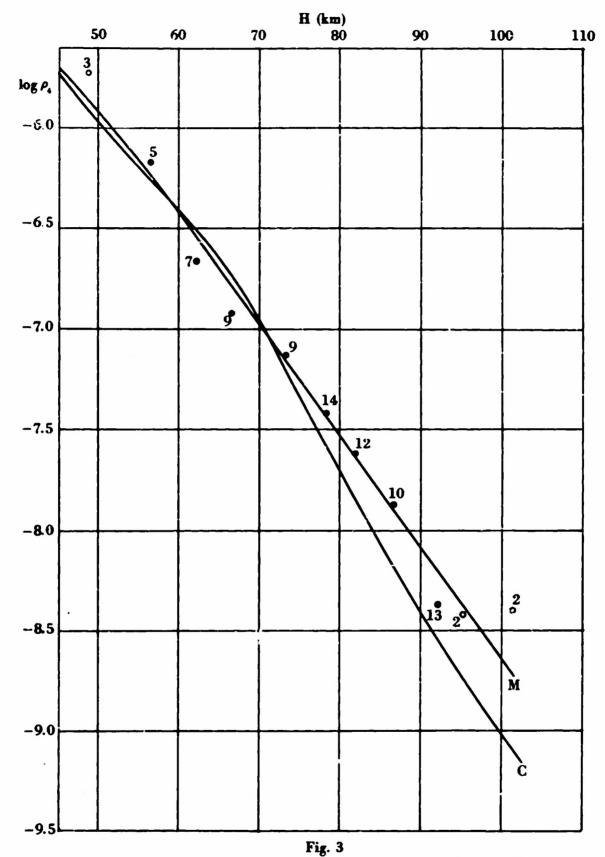
TABLE X

(Weighted Means within 5-Km Groups; Geminid Densities Corrected)

H (1cm)	$\log \rho_1$	<u>م</u>	log P2	Δ2	log ρ_{3}	Δ,	Δ',	log $\rho_{_{k}}$	Δ,	Δ',	•
48.83	-6.12	-0.20	-6.10	-0.18	-5.75	+0.17	+0.10	-5.73	+0.19	+0.12	3
56.63	-6.48	-0.22	-6.47	-0.21	-6.22	+0.04	+0.01	-6.17	+0.09	+0.06	5
62.32	-6.76	-0.25	-6.88	-0.37	-6.71	-0.20	-0.16	-6.66	-0.15	-0.11	7
66.63	-7.02	-0.28	-7.07	-0.33	-6.96	-0.22	-0.17	-6.92	-0.18	-0.13	9
73.43	-7.20	0.00	-7.20	0.00	-7.13	+0.07	+0.03	-7.13	+0.07	+0.03	9
78.30	-7.36	+0.21	-7.39	+0.18	-7.36	+0.21	+0.06	-7.42	+0.15	0.00	14
81.98	-7.63	+0.22	-7.57	+0.28	-7.60	•0.25	+0.02	-7.62	+0.23	0.00	12
86.8	-7.96	+0.23	-7.78	+0.41	-7.86	+0.33	+0.04	-7.87	+0.32	+0.03	10
92.2	-8.35	+0.19	-8.07	+0.47	-8.23	+0.31	-0.03	-8.37	+0.17	-0.17	13
95.2	-8.46	+0.27	-8.12	+0.61	-8.30	+0.43	+0.07	-8.42	+0.31	-0.05	2
101.3	-8.47	+0.63	-8.12	+0.98	-8.30	+0.80	+0.41	-8.40	+0.70	+0.31	2

 $\Delta_1, \, \Delta_2, \, \Delta_3, \, \Delta_4$ are residuals of the respective log $\rho's$ from the Rocket Profile C .

 Δ_3' and Δ_4' are residuals of $\log \rho_3$ and $\log \rho_4$ from the Massachusetto Profile H .



Corrected Mean Atmospheric Densities ρ_{η} from Massachusetts Meteors. The figures which eccompany the plotted points indicate the number of observations used to compute the mean. Means from fewer than five observations are shown as open circles. The rocket profile C and the new Massachusetts profile M are shown for communison.

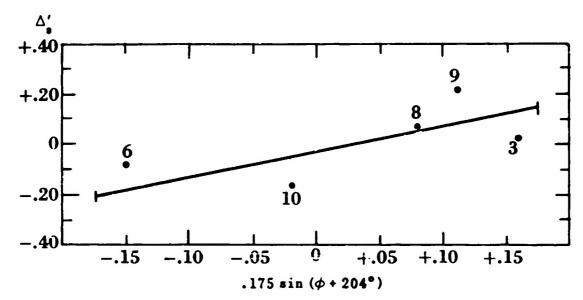


Fig. 4

Seasonal Effect in Massachusetts. Means of the residuals Δ_3' of $\log \rho_3$ from profile M for heights between 65 and 85 km. The number of residuals which concurred to form the individual means is indicated in the diagram. The straight line represents the result of the least-squares solution.

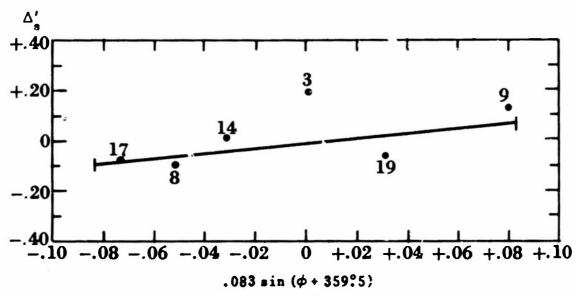


Fig. 5

Seasonal Effect in New Mexico. Means of the residuals Δ_3 of $\log \rho_3$ from the rocket profile C for heights between 65 and 85 km. The number of residuals which concurred to form the individual seans is indicated in the diagram. The straight line represents the result of the least-squares solution.

As we have previously mentioned, the distribution of Massachusetts meteors with regard to seasons, velocities and heights is quite bad. A sizable fraction of New Mexico meteors yielded more than one deceleration in the course of the visible trajectory. In Massachusetts with smaller cameras, poorer sky and slower, two-blade shutters, multiple-deceleration meteors were the exception rather than the rule. The result is that even in the region between the heights of 65 and 85 km there is little overlap in heights and velocitien, and the situation is made worse by the preference—fortuitous or not—of high-velocity meteors for the summer season.

As a consequence of this situation, the parameters of the equation of condition (5) are atrongly interrelated and the least-squares solutions computed on its basis lose much of their significance. This must be kept in mind when results from this analysis are compared with

previous results.

The instability of the solution is made quite evident by an inspection of the d column in Table VIII. While for New Mexico the value of d was quite stable in the vicinity of -1.1 for all comparison profiles and changes of parameters, here it changes very drastically with the slope of the comparison profile. If we make e=0, i.e., if we do not allow for an error in the slope of the comparison profile, the strong dependence of velocities on heights will cause d to assume practically any value, and even to change sign, according to the profile which is taken for comparison. The comparison profile has also a strong influence on the amplitude of the seasonal effect — although much less on its phase.

Of all the solutions, No. 2 is the one which can be most directly compared with the results of Technical Report No. 2, Eq. (14). The comparison profile was in one case the N.A.C.A. profile⁽⁷⁾, in the other profile A, but due to the presence of the eH term in the equation of condition, the conditions are quite similar in both cases. The value of d was -0.91 in the old, -0.76 in the new solution.

After this digression, we can go back to the outlined reduction of the Massachusetts ρ_1 's. Table IVs gives $\log \rho_2$ and $\log \rho_3$ as computed from ρ_1 , with their residuals Δ_2 and Δ_3 from profile C. Mean values of $\log \rho_3$ taken in 5-km intervals were plotted against H and a smooth curve drawn through the points. The residuals from this curve, in the height range between 65 and 85 km (Geminids excluded) were analyzed for seasonal effect and the resulting seasonal correction applied to ρ_3 . A new curve (almost identical with the preceding one) was drawn through these corrected densities and was assumed to be the final Massachusetts density profile (M in Table VI). The residuals of $\log \rho_2$ from this profile are designated Δ_3 in Tables IVs and IVb. A final analysis of Δ_3 in the same interval as before, gives

$$\Delta'_3 = -0.033 + 0.175 \sin (\phi + 204^\circ)$$
 $\pm .032 \pm .040 \pm 16$ (12)

The final corrected densities ρ_{ij} [i.e., $\log \rho_{ij}$ with the correction -0.175 Lin (ϕ + 204°) applied to them, and the extra correction +0.10 applied to Geminida] are shown in Tables IVs and IVb. Means in 5-km height intervals are given in Table X and plotted in Fig. 3.

Report/No. 2. There the fluctuation was given as a function of the mean normal temperature T in Boston, with a coefficient of +0.015 per °C. This would correspond to a semi-amplitude of 0.18 with a minimum toward the end of January. Here the semi-amplitude is 0.175 with a minimum on March 9. It is perhaps significant that, while the amplitude of the seasonal fluctuation turns out to be twice as large as in New Mexico, its probable error is just about the same in both localities. Although some caution should be exercised in interpreting this result in view of the bad distribution of the meteors in Massachusetts, it might in licate a change of the seasonal fluctuation with latitude; if so, the change is definitely in the expected direction.

The final, corrected densities for Massachusetts seem to be in fair agreement with the New Mexico profile up to 75 km. For greater heights, however, the Massachusetts densities seem to be systematically higher. It would be highly suggestive to attribute this divergence to a latitude effect. We hope that future observations will throw more light on this important point.

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